

Constraining uncertainty of *in situ* chamber-based estimates of the stable carbon isotope ratio of soil-respired CO₂ via advances in automated sampling system technology

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Introduction

- Chamber based measurements of soil CO₂ flux are a key tool in understanding the role of soil in ecosystem carbon cycling.
- Estimation of stable carbon isotope ratio ($\delta^{13}\text{C}$) of soil-respired CO₂ via the Keeling mixing model provides a more comprehensive measurement which can help disentangle the drivers of CO₂ fluxes.
- Commercial enhanced cavity absorption gas analysers utilise H₂O corrections which can create uncertainty for measurement of low abundance gases such as ¹³CO₂ during large, rapid changes in H₂O concentration. They also exhibit a non-negligible CO₂ concentration dependency in the $\delta^{13}\text{C}$ measurement.
- H₂O transients during chamber-based soil flux measurements can be very large, whilst CO₂ transients are smaller but often significant. This has traditionally made *in situ* measurement of $\delta^{13}\text{C}$ challenging, where the use of powered solutions or consumable chemicals can be problematic. Presented here is a passive solution to address these issues.

Sensitivity of the Keeling method to CO₂ and H₂O concentration changes

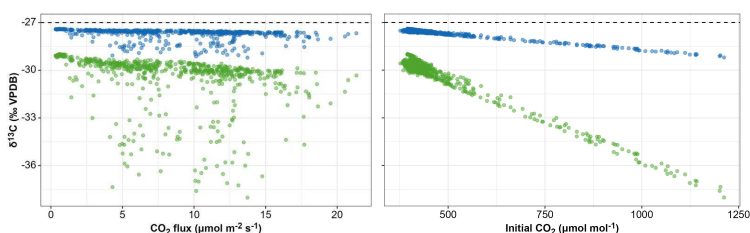


Figure 1: Synthetic dataset with introduced delta dependency on total CO₂ concentration at two magnitudes noted in the literature 0.5 ‰ $\mu\text{mol mol}^{-1}$ reported in Tazoe *et al.*, 2011 (green) and 0.1 ‰ $\mu\text{mol mol}^{-1}$ reported in Braendholdt *et al.*, 2019 (blue). The synthetic data $\delta^{13}\text{C} = -27$ (dashed line).

Fluxes were generated using the flux fit parameters from 655 measurements in Lincoln, NE. The Keeling model estimated the $\delta^{13}\text{C}$ of respired CO₂, but its regressions are sensitive to errors in initial CO₂ and concentration range. Delta dependence to CO₂ range (left) is limited due to the typically narrow range during flux measurements. However, high initial CO₂ concentrations (right) can cause significant underestimation of $\delta^{13}\text{C}$ source values.

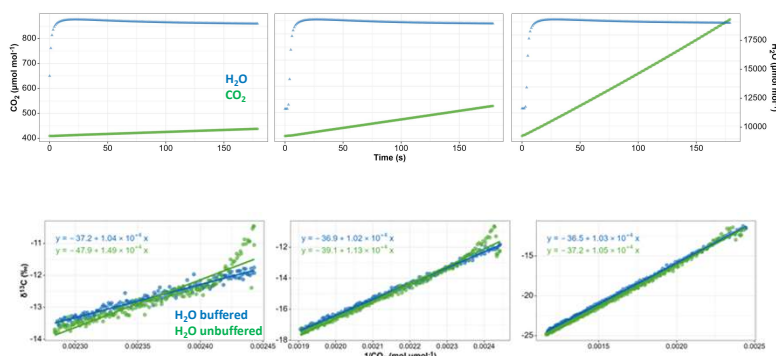


Figure 2: A range of H₂O and CO₂ transients were generated using the LI-6800 photosynthesis system to simulate common concentration changes observed within a soil flux measurement (top). Pure CO₂ was injected at $\delta^{13}\text{C} = -36.7\text{‰}$ and the system was verified to not cause CO₂ isotopic fractionation.

Keeling regressions (bottom) show improved source estimation accuracy with H₂O buffering (blue) vs. unbuffered (green), using the system shown in Figure 4. Across 18 simulations, mean source estimates were -36.8‰ (buffered) and -37.1‰ (unbuffered), with standard deviations of 0.42 and 3.78, respectively. This corresponds to mean inaccuracies of 0.79% (buffered) and 6.26% (unbuffered). Errors increased with smaller CO₂ ranges due to greater noise sensitivity.

Reducing $\delta^{13}\text{C}$ sensitivity to CO₂ concentration

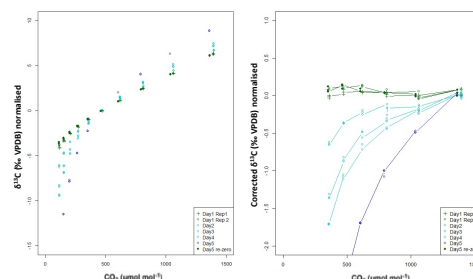


Figure 3: CO₂ was injected from a tank of pure CO₂ at a range of concentrations controlled by the LI-6800 photosynthesis system without correction (1). The CO₂ mixing system was verified to not cause fractionation. The experiment was repeated over a series of days (2) and data corrected for CO₂ concentration. $\delta^{13}\text{C}$ shows daily drift but re-zeroing collapses CO₂ concentration sensitivity to its 'at factory' value and can be corrected for.

Automated field measurement system

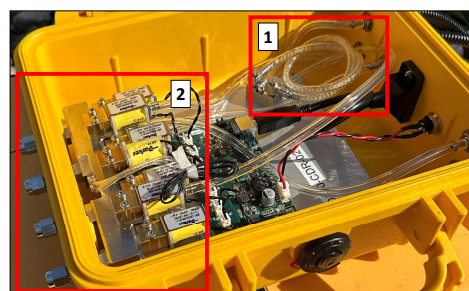


Figure 4: Air is passed through the LI-8250 calibration manifold passive water transient buffer (1) before the Trace Gas Analyser. Separately the calibration manifold provides connections for up to four gas tanks allowing for automated baselining in the field (2) addressing the delta dependence on CO₂ mole fraction shown in Figure 3.

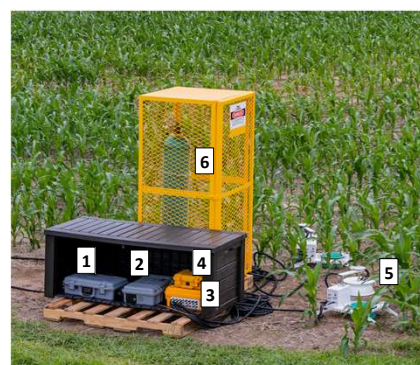


Figure 5: Photograph of field deployed measurement system

- LI-7825 CO₂ Isotope / NH₃ Trace Gas Analyser
- LI-7820 N₂O / H₂O Trace Gas Analyser
- LI-8250 Multiplexer
- LI-8250-02 Calibration Manifold
- 8200-104 Long-Term Chamber
- Zero gas

Conclusions

- Rapid changes in H₂O concentration represent a significant challenge in accurate isotopic flux quantification.
- The passive solution presented here offers control of the water transient across the duration of soil flux measurements without the need for powered drying or consumable chemicals and demonstrated large improvements in source composition estimation.
- Instruments show sensitivity to CO₂ concentration; however daily zeroing ensures that sensitivity returns to a known state and can be calibrated for.
- The combination H₂O buffering and CO₂ baselining allows for automated measurement of $\delta^{13}\text{C}$ in the field with low power requirements.

